Crystal-chemical transformations and iron reduction in clay minerals of two Brazilian mangrove soils

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The formation of clay minerals in mangrove soils is an important key to understand pedogenic and biogeochemical processes. To investigate the chemical transformations and the role of Fe reduction in clays, two Brazilian mangroves were selected (Bragança, Pará State, BR - humid Amazonian coast; Acaraú, Ceará State, AC - semiarid northeastern coast). Samples of 2-0.2 μm and < 0.2 μm size fractions were studied using modeling of XRD patterns (program ClaySIM[®]), X-ray fluorescence spectroscopy (XRF) and Mössbauer spectroscopy.

Table 1 – Phase distribution and crystal-chemical parameters used for XRD modeling procedure (corresponding experimental and calculated XRD patterns are shown in Fig. 1)

% layers Fe kaol Fe sm Fe il K il % phase Components Sample Nave Nmax





	2-0.2 μm size fraction												
	BR1_20-30	73	K-S	98-2	0.71	1.40			58	42			
		1	K-S	90-10	0.71	1.40			10	3			
		11	I-S	90-10		1.40	0.45	0.50	7	1			
		15	Illite	100			0.45	0.50	25	15			
	AC1_0-5	11	K-S	98.5-1.5	0.54	1.40			40	25			
		3	K-S	90-10	0.54	1.40			12	7			
		18	Smectite	100		1.40			7	1			
		1	Vermiculite	100		0.98			12	6			
		60	Fe-illite	100			1.40	0.50	22	16			
		8	Al-illite	100			0.20	0.80	30	20			
< 0.2 μ m size fraction													
	BR1_20-30	28	K-S	99-1	0.00	1.10			35	25			
		14	K-S	85-15	0.00	1.10			12	7			
		50	K-S	10-90	0.00	1.10			7	1			



Fig. 1 – Examples of XRD patterns from the two mangroves. Black lines – experimental glycolated patterns. Gray line – fitted calculated patterns, using the program ClaySIM[®]

Fig. 2 – Examples of Mössbauer spectra of samples from BR mangrove soil. A- Sample from 2-0.2 μ m size fraction. B – Sample from < 0.2 μ m size fraction. Circles – experimental data; Black line – overall fit; Gray line – octahedral Fe³⁺ doublet; Dash line – octahedral Fe²⁺ doublet

The deposited kaolinite is progressively transformed into smectite via K-S. This step includes the formation of a new tetrahedral sheet, gradual substitution of cations in tetrahedral (Al for Si) and octahedral sites (Fe+Mg for Al), and intercalation of cations, leading to the formation of Fe-rich smectite layers. The second step is the illitization of smectite layers, which includes further Fe incorporation and reduction. This is concluded because both total Fe and Fe²⁺/TotFe proportion in the 2-0.2 μ m size fraction correlated very well with Fe-illite contents from XRD. The AC samples have higher Fe-illite content because continental 2:1 clays are deposited in the mangrove and they are transformed more quickly into Fe-illite than kaolinite. We conclude that Fe sequestration and reduction during illitization are important biogeochemical processes operating in mangrove soils.

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